

# Course Remote Sensing and Climate

**Summer 2026**  
**University of Frankfurt**

## **2. Day (R. Hollmann)**

- Content and Structure of Course
- Motivation and background
- **A short introduction in Physics**
- Radiative Transport and retrieval basics
- Satellite orbits and instruments
- Climatologies based on satellite instrument and usage



## **Notes to PDF-Version**

**This PDF is based on the presentation given in the summer term 2025 of University Frankfurt as part of the course „remote sensing and climate“.**

**It is meant to recall the content the content of the lecture and is for the personal use of the participants of the course and not for a wider audience.**

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# A short introduction in Physics

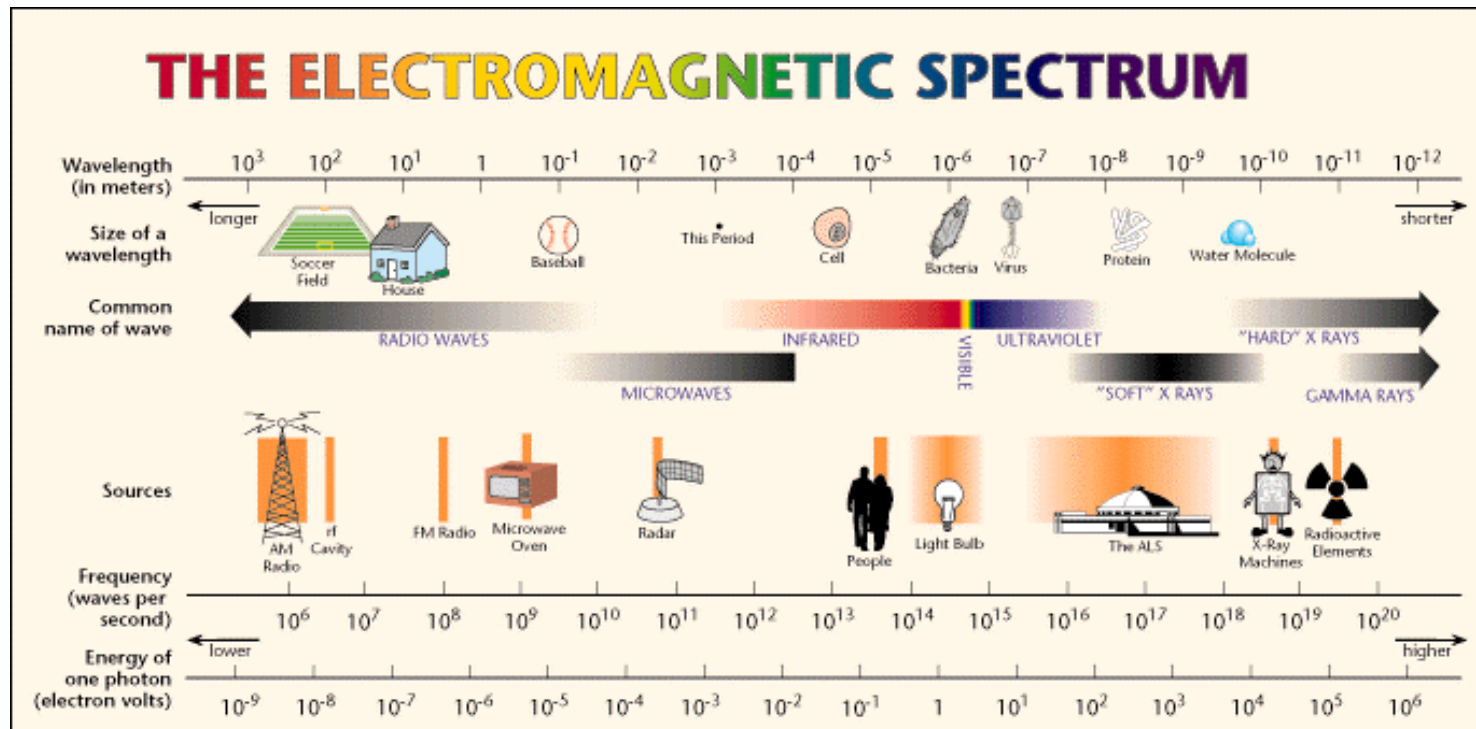
## Content

- electromagnetic spectrum
- Solid angle, Radiance, Flux
- Black body and Planck function
- Reflection, Emission

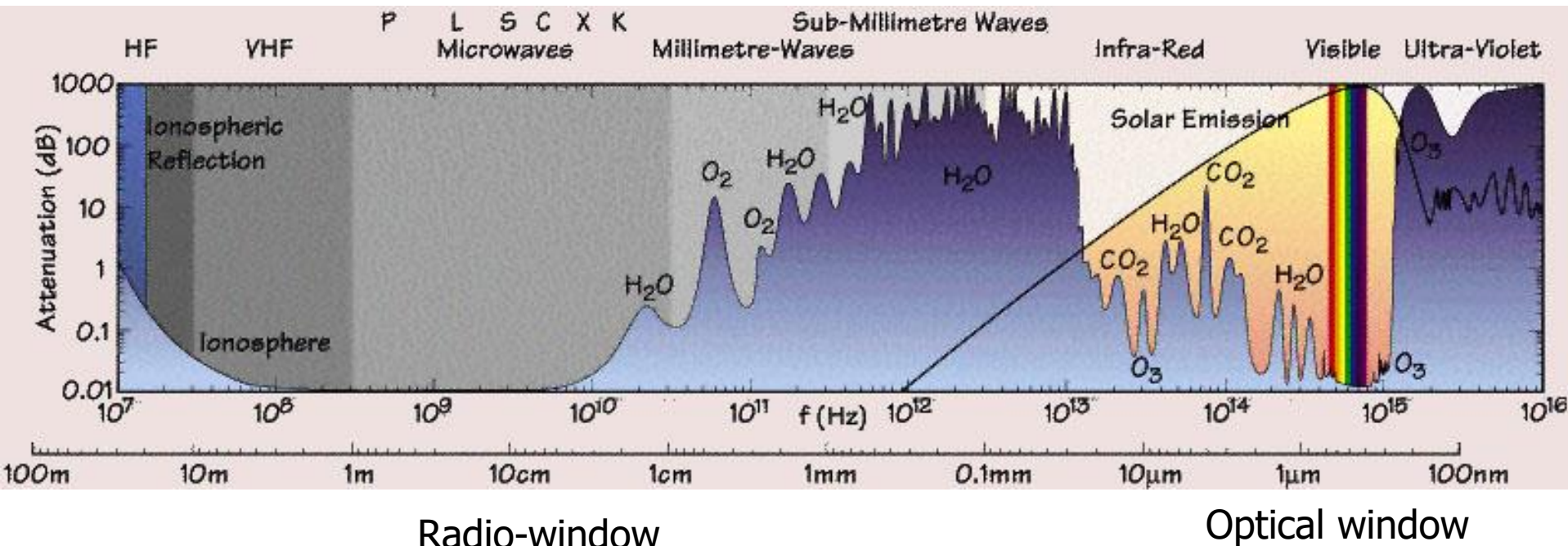


# Electromagnetic Spectrum (I)

- Electromagnetic radiation is a form of energy propagation. It can be understood as a wave, i.e. as a periodically changing electromagnetic field. It is characterized by the frequency or wavelength.
- The entirety of the wavelengths occurring in electromagnetic radiation is represented in the electromagnetic spectrum:



# Electromagnetic Spectrum (II)

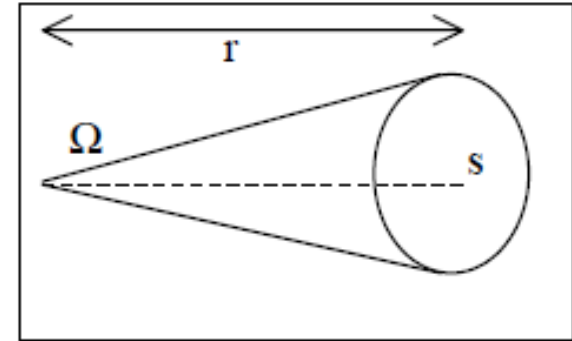


Which parts are often used in satellite remote sensing?

# Solid Angle $\Omega$

**Solid Angle** (often Steradian) is defined as ratio of an area  $s$  of a cone to the radius  $r$  of this cone:

$$\Omega = s / r^2.$$



Expressed as differential  $d\Omega$  and using

$$ds = (r d\theta) (r \sin(\theta) d\phi)$$

(sphere co-ordinates!) we yield :

$$d\Omega = \frac{ds}{r^2} = \sin(\theta) d\theta d\phi$$



# Radiance

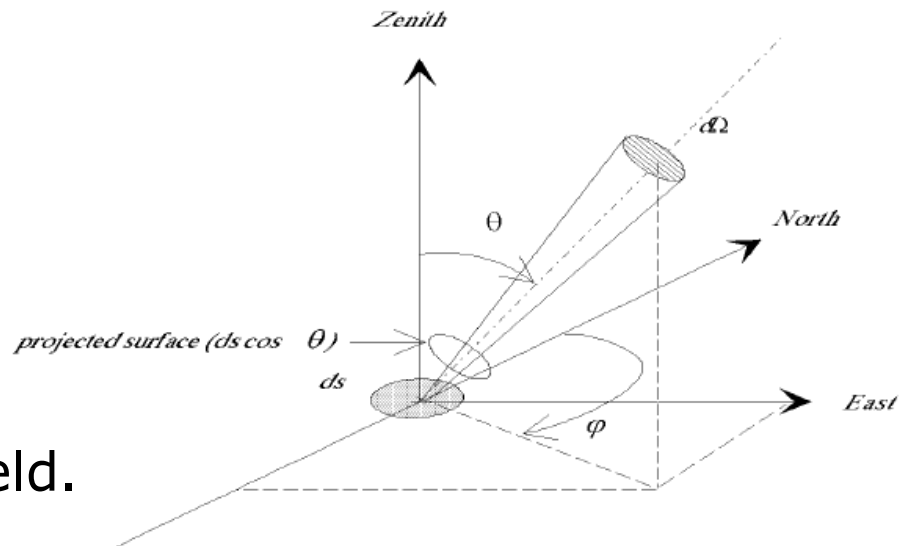
The **radiance L** or intensity (I) is the radiant energy in one direction per unit of time, wavelength and solid angle (perpendicular to the axis) and is defined as:

$$L_{\lambda} = \frac{dE_{\lambda}}{ds \cos(\theta) d\Omega dt d\lambda}$$

Unit: W / (m<sup>2</sup> sr μm).

L (or I) is monochromatic

The radiance as a function of direction enables a complete description of the radiation field.



# Characteristics of radiance

In general,  $L$  is a function of

- direction
- location
- wavelength
- time

= 7 independent variables

The integration of  $L$  over solid angle provides the radiation flux density.

**Special cases:**

**$L$  independent of direction:**

**ISOTROP**

**$L$  independent of location:**

**HOMOGEN**





# radiant flux and radiance flux density $\Phi$

The **radiant flux  $\Phi$**  describes the radiant energy that is transported, received or emitted per unit of time. It provides neither information about the directional distribution nor about the spatial distribution of the radiation.

$$\Phi_{\lambda} = \frac{dE_{\lambda}}{dt d\lambda}$$

Unit is Watt.

The **radiance flux density or irradiance** describes the radiation flux through a certain area (dA) and thus provides information about the area distribution of the radiation:

$$F_{\lambda} = \int_{\Omega} I_{\lambda} \cos(\theta) d\Omega = \frac{dE_{\lambda}}{ds dt d\lambda} = \frac{d\Phi_{\lambda}}{ds}$$

Unit is Watt / (m<sup>2</sup> μm).



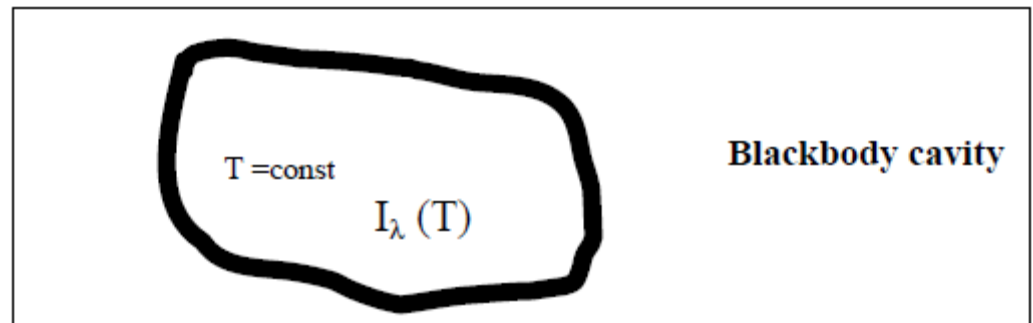
# Blackbody and Thermodynamic equilibrium

A **blackbody** is a body that absorbs all incident radiation.

**Thermodynamic equilibrium** describes the state of radiation and the body within an isolated environment at constant temperature.

## Blackbody temperature

The blackbody radiation is the radiation field within a body in thermodynamic equilibrium.



# Characteristics of blackbody radiation

The **emitted radiation** of a blackbody is

- **Isotropic**
- **homogeneous**
- **not polarized**
- **depends only on the temperature and wavelength**

Two **blackbodies** with the **same temperature** emit exactly the **same radiation**

A **blackbody** emits **more radiation** than **all other objects / types** at the **same temperature**

The **Planck function** describes the **radiance** of the **blackbody**.

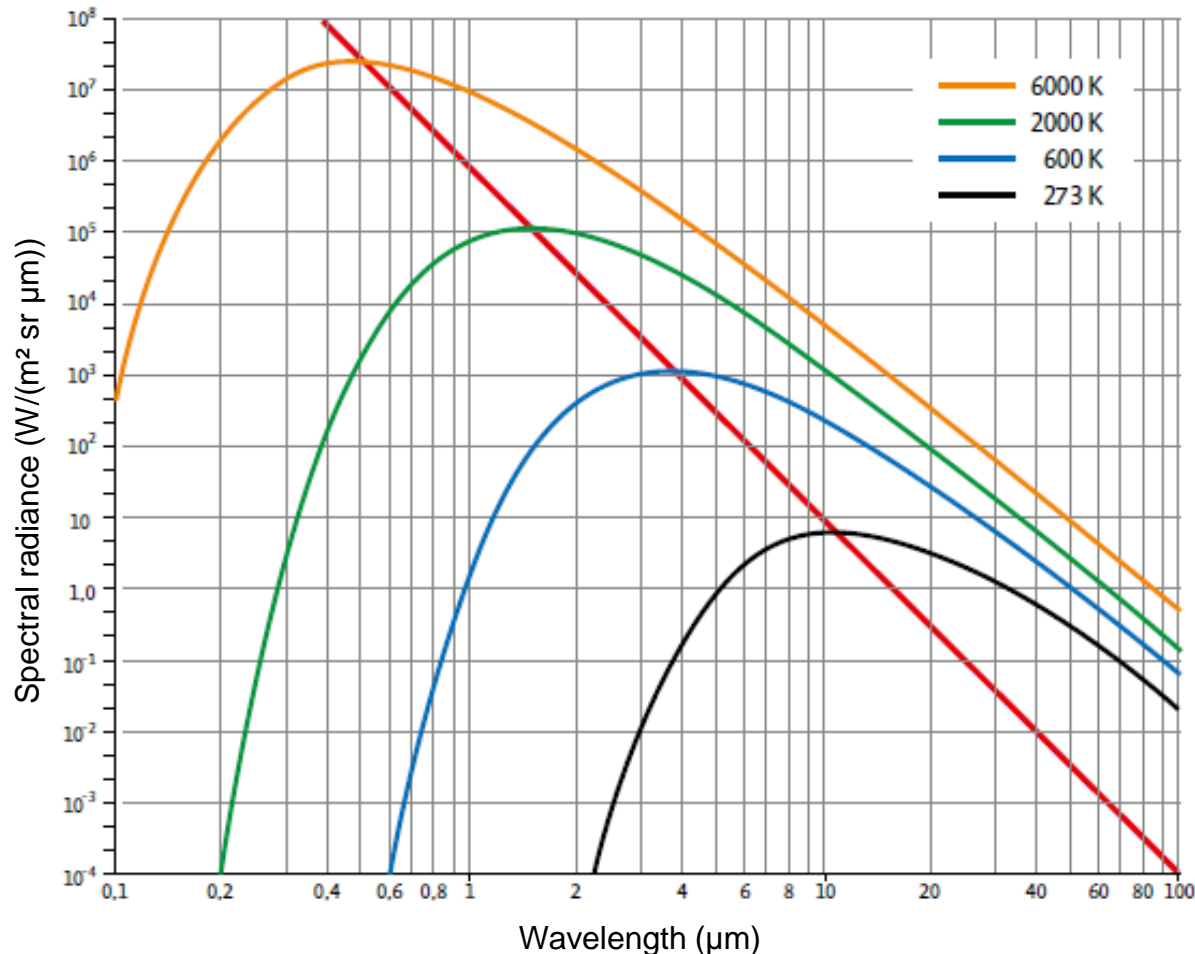
$$B_{\lambda}(T) = \frac{2hc^2}{\lambda^5 (\exp(hc / k_B T \lambda) - 1)}$$



# Planckfunction

What about the Planckfunction for the Earth?

What about the Planckfunction for the Sun?



Planckfunction  
for different  
temperatures

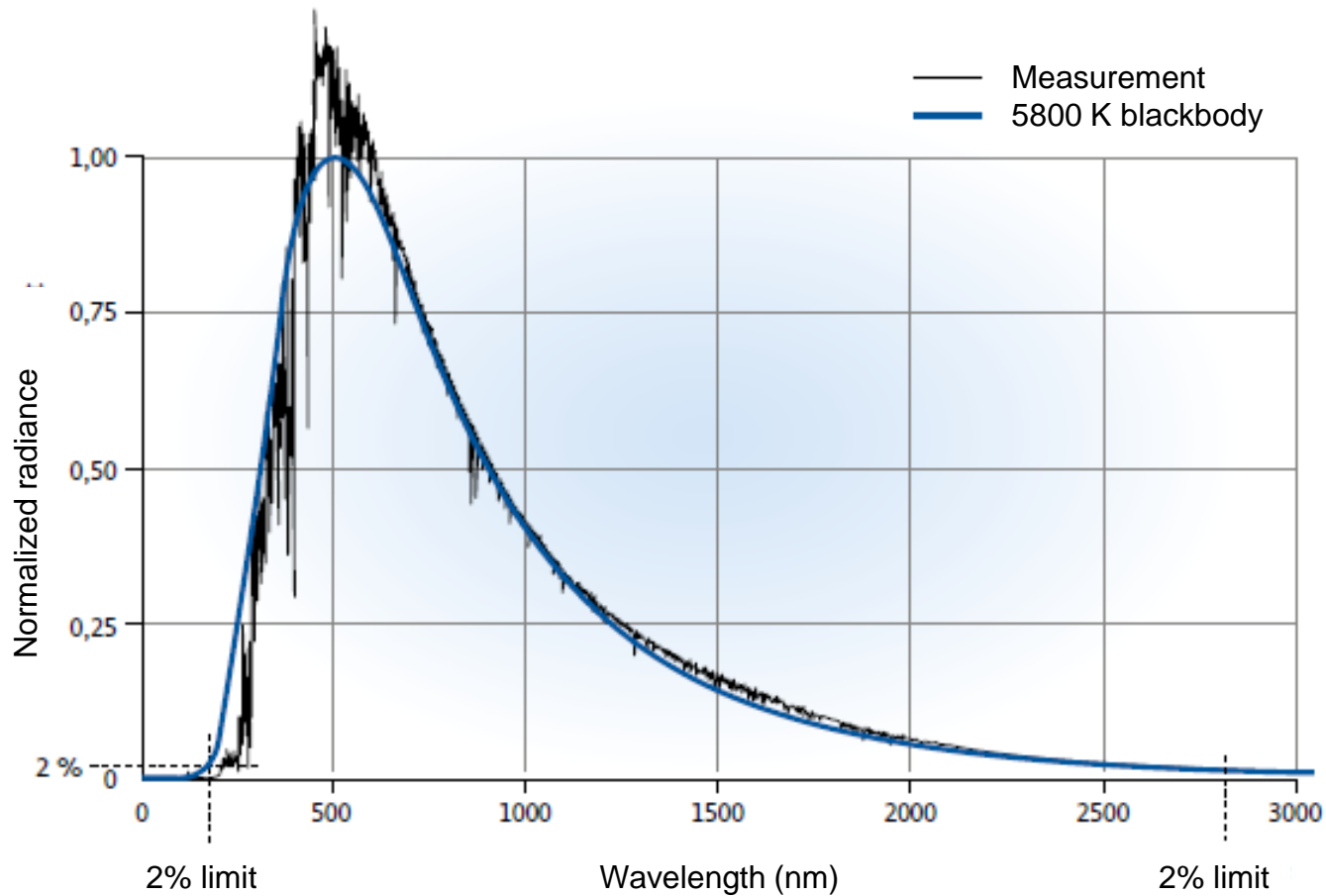
Spectral radiance of  
blackbodies with  
6000, 2000, 600 and  
273 K.

Red line indicates the  
respective radiation  
maximum.

Köpcke et al. (2012),  
adapted



## What about the Planckfunction for the Sun?



Köpcke et al. (2012)



# Planckfunction

$$B_{\lambda}(T) = \frac{2hc^2}{\lambda^5 (\exp(hc / k_B T \lambda) - 1)}$$

For large Wavelengths  
(Rayleigh-Jeans Law):  $\lambda \rightarrow \infty \Rightarrow B_{\lambda}(T) = \frac{2k_B c}{\lambda^4} T$

Linear Approximation is important for passive microwave

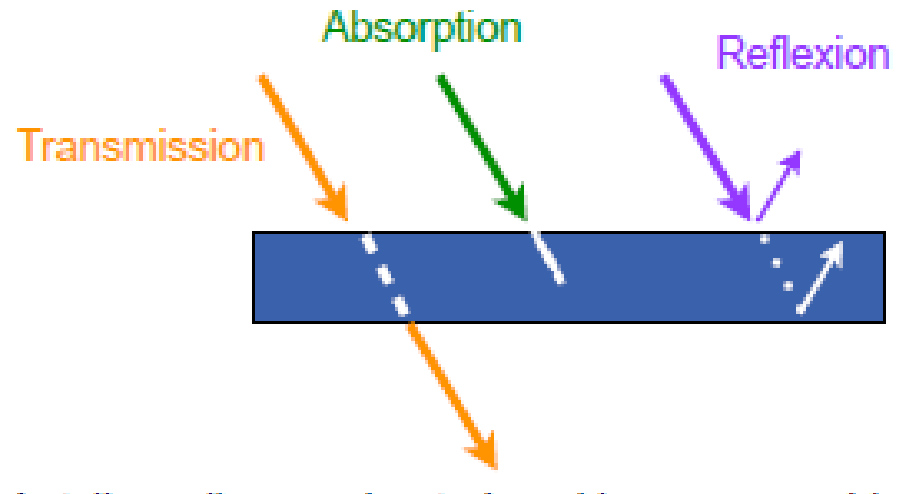
Stefan Boltzmann Law:  $F = \sigma T_B^4, \quad \sigma = 5.67 \cdot 10^{-8} \text{ W } /(\text{m}^2 \text{K}^4)$



# Interaction of electromagnetic radiation

Part of the electromagnetic radiation hitting a body is **reflected**, another part is **absorbed** and the rest penetrates the body which is called **transmission**.

The **individual proportions** vary greatly and **depend** on both the **characteristics** of the body and the **wavelength** of the radiation in question.



# Extinction and Emission

Electromagnetic radiation **in the atmosphere** is influenced by gases, aerosols and cloud particles:

**extinction** (“attenuation”) leads to a reduction in radiance

**absorption**: Transformation of electromagnetic energy into other energy

**scattering**: Redistribution of electromagnetic energy in a different direction

(a) elastic -> constant wavelength

(b) inelastic (Raman) -> wavelength changes

**Emission** leads to an increase in radiance

- essentially in the thermal range





# Emission and Reflection of surfaces

**Energy conservation:** Absorption + Transmission + Reflection = 1

## **Emissivity of surface:**

Generally, depends on

- Direction
- surface temperature
- Wavelength
- Surface properties

**Thermal IR** ( $4\mu\text{m} < \lambda < 100\mu\text{m}$ ) all surface with an emission larger than 0.8 are nearly independent of direction :

$$L_{\lambda} = \varepsilon_{\lambda} B_{\lambda}(T_s), \quad F = \varepsilon \sigma T_s^4$$

**Solar:** ( $0.1 \mu\text{m} < \lambda < 4 \mu\text{m}$ ), emission negligibly small

**Microwave** ( $0.1 \text{ cm} < \lambda < 100 \text{ cm}$ ), emission depends strongly on type and surface state (e.g. ice/water)



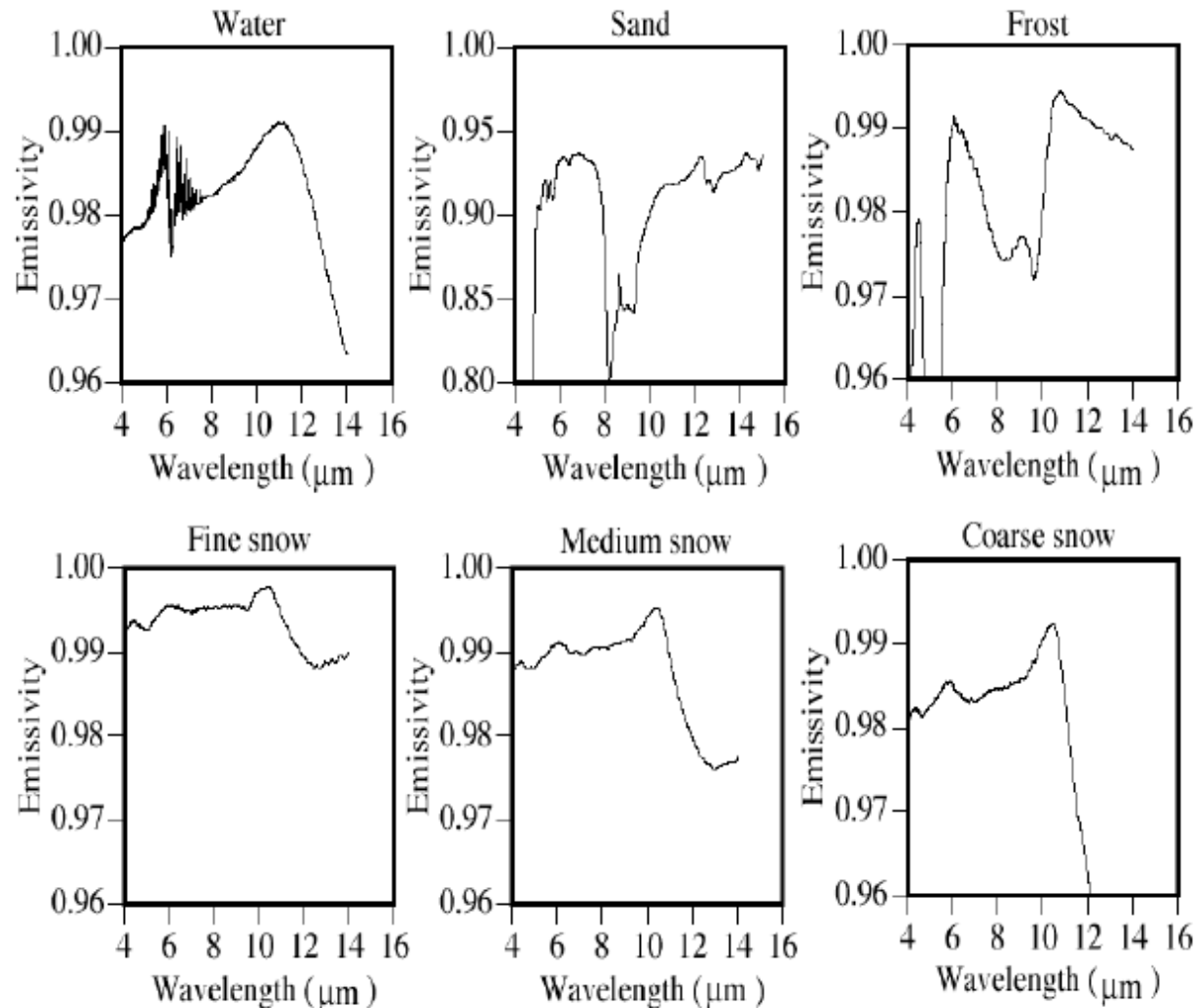
# Emissivity of surface types (3)

Surface	Emissivity
Water	0.993-0.998
Ice	0.98
Green grass	0.975-0.986
Sand	0.949-0.962
Snow	0.969-0.997
Granite	0.898

Emissivity of surfaces types (10-12  $\mu\text{m}$  domain)



# Emissivity of surface types (2)



Spectral emissivity of surface types



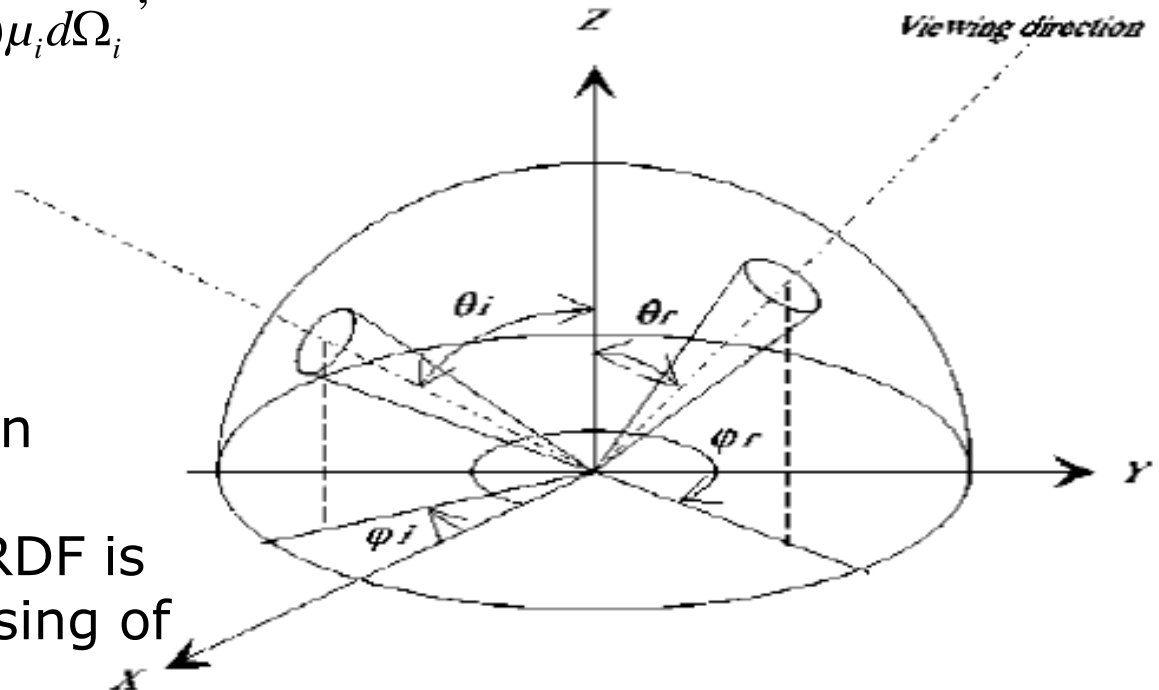
# Reflection of Surfaces

Surfaces are anisotropic, which is why a **BRDF (bi-directional reflectance distribution)** is introduced. It characterizes the angular dependence of the ground reflection:

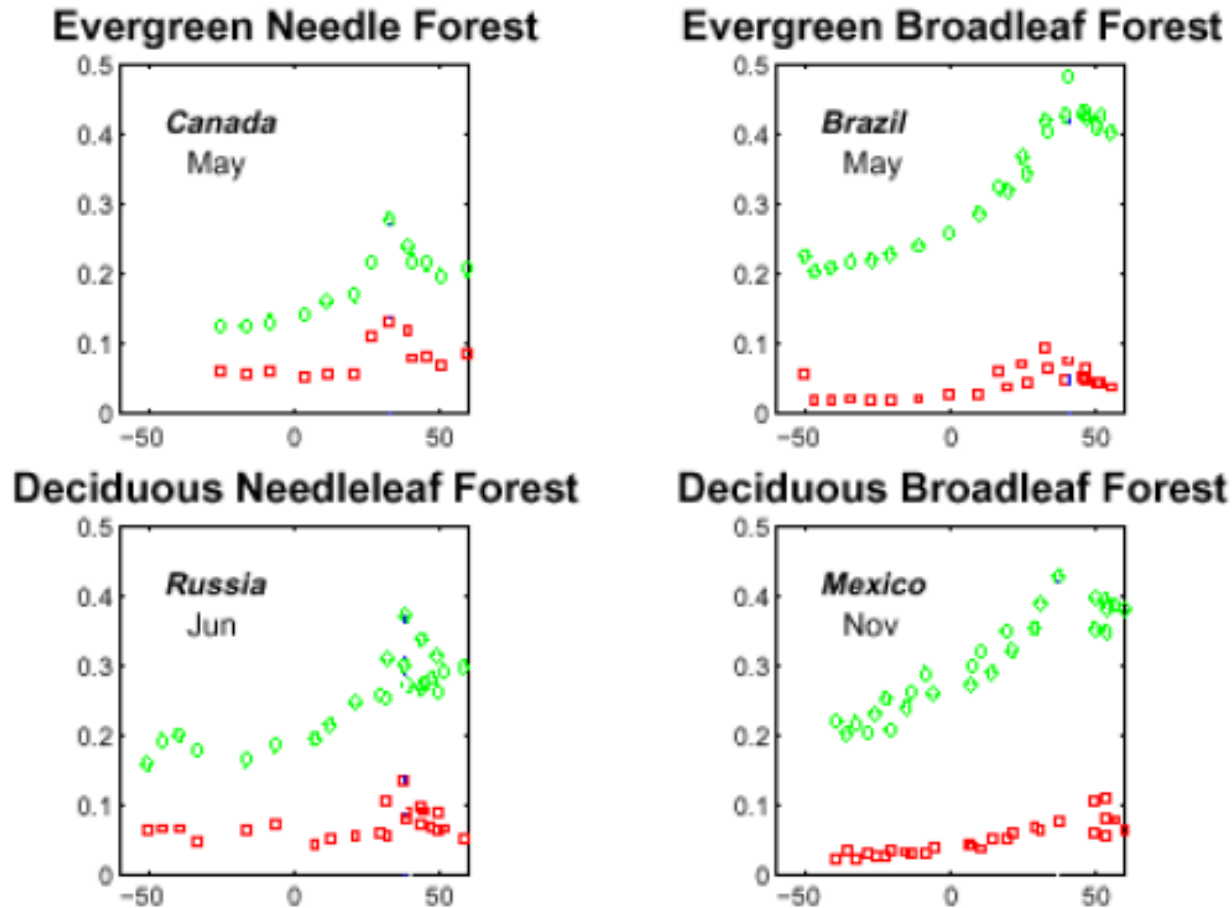
$$R(\mu_i, \varphi_i, \mu_r, \varphi_r) = \frac{\pi dI^\uparrow(\mu_r, \varphi_r)}{dI^\downarrow(\mu_i, \varphi_i) \mu_i d\Omega_i},$$

$$\mu_i = \cos(\theta_i), \mu_r = \cos(\theta_r)$$

- the BRDF depends on wavelength.
- Knowledge of the BRDF is needed in remote sensing of land surfaces



# BRDF



Example of the BRDF for 4 land surface types.  
For two different visible channels 0.6  $\mu\text{m}$  (green), 1.6  $\mu\text{m}$  (red).

